

Spatial evolution of tidal spectra: From shelf seas to estuaries

Emil Stanev

with contributions of

*J.Y. Zhang, G. Grashorn, J. Pein, B. Jacob, R. Al Nadhairi, J. Staneva, J. Schulz-Stellenfleth,
and A. Valle-Levinson*

*International Coordination Workshop (ICW4) of the
Coastal and Shelf Seas Task Team (COSS-TT)
&
ARCOM Pilot Workshop
31 August – 4 September 2015,
Instituto Superior Técnico, Lisbon, Portugal*

Motivation:

Mesoscale and submesoscale dynamics

Dynamics of spectral composition, tidal wave transformations and generation of overtides

Address open sea and estuarine dynamics in one study.

Tidal spectroscopy of shallow coastal areas and nonlinear distortion and friction: Munk and Cartwright (1966), Gallagher and Munk (1971), Ianniello (1977), Jay and Smith (1990), Le Provost (1991), Parker (1991).

*In coastal regions, differently from the open ocean, the tidal spectrum is more complex (**overtides and compound tides**).*

*Nonlinear advection terms are responsible for the generation of even harmonic overtides (**M4**).*

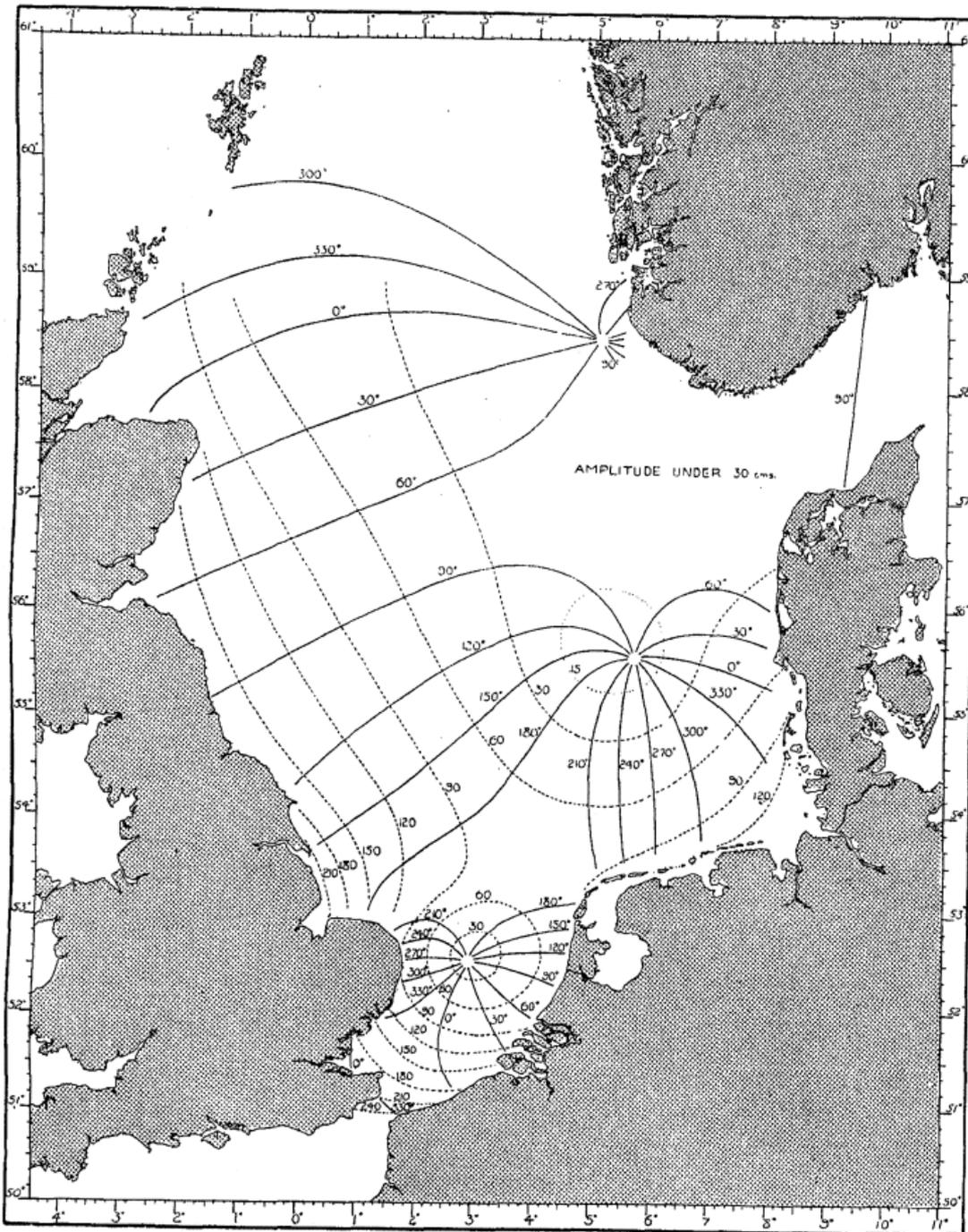
*Nonlinear friction is responsible for producing odd harmonic overtides (**M6**).*

Past works have overlooked:

- (1) the horizontal patterns of overtides at the transition between the coastal zone and the open ocean,*
- (2) the role of bathymetric channels,*
- (3) the role of baroclinicity,*
- (4) the seamless approach.*

*One basic **practical problem** is the assimilation of altimetry into numerical tidal models (Le Provost et al. 1994; Le Provost 2001).*

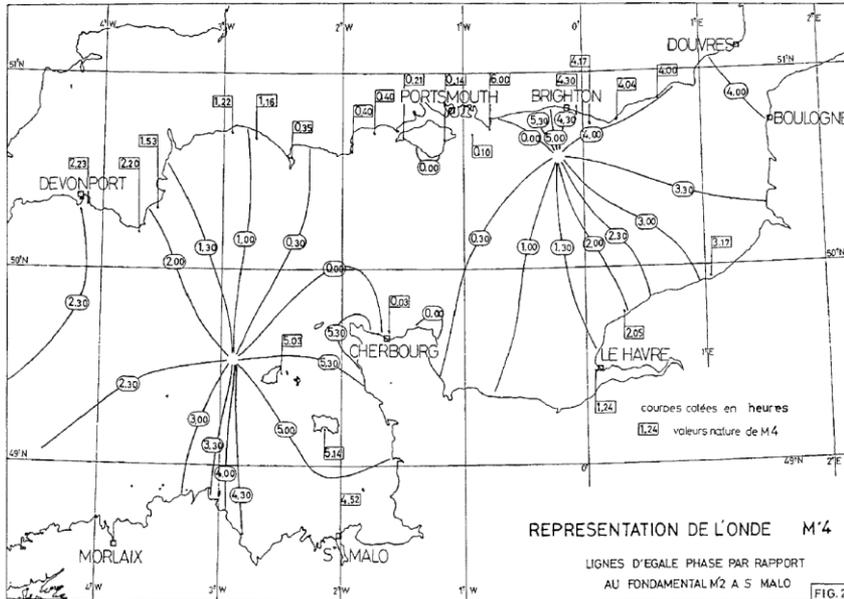
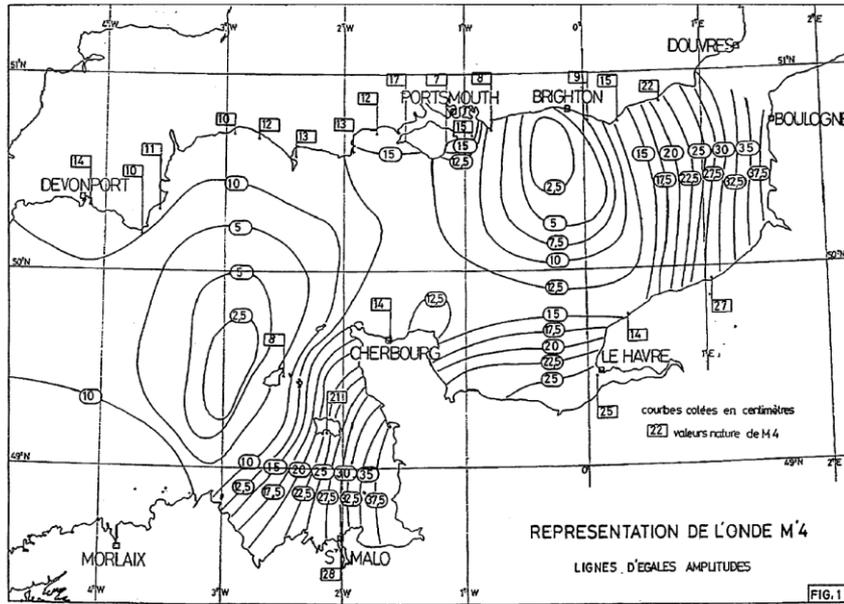
1924



Proudman and Doodson: How the fundamental dynamical equations of the tides may be used to obtain a knowledge of the distribution of surface elevation from observations.

1970

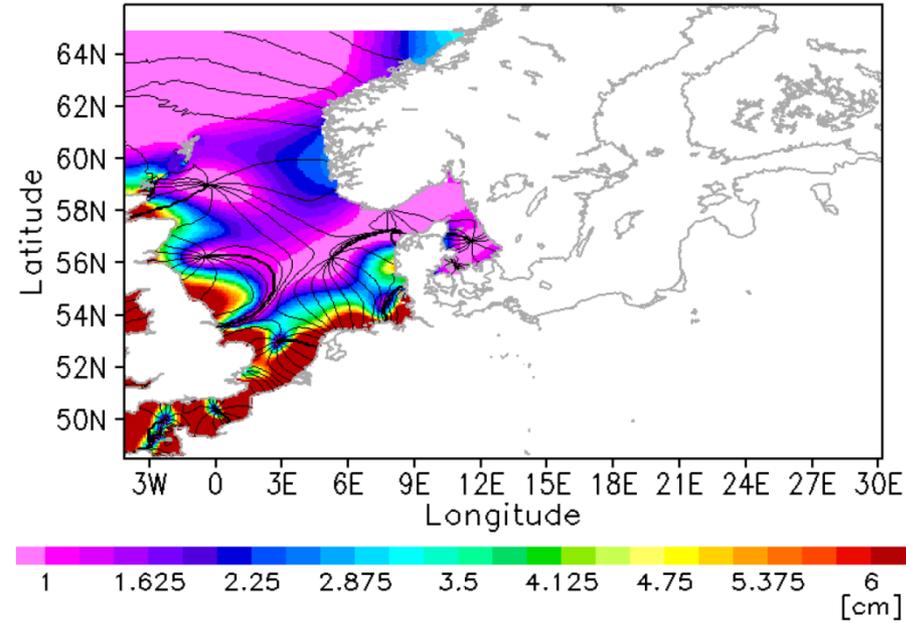
The first cotidal chart of M4 in this region



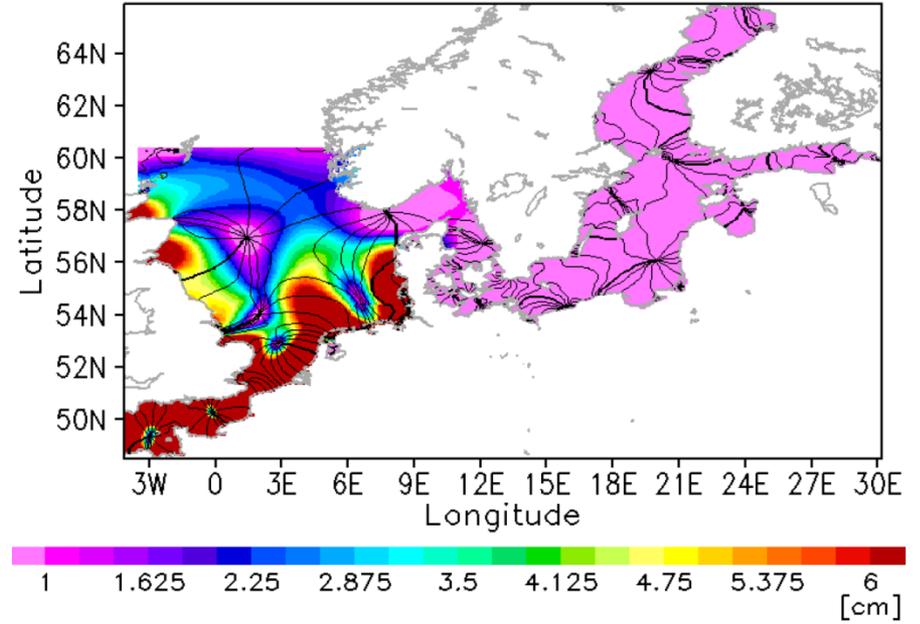
Chabert d'Hières and Le Provost

2015

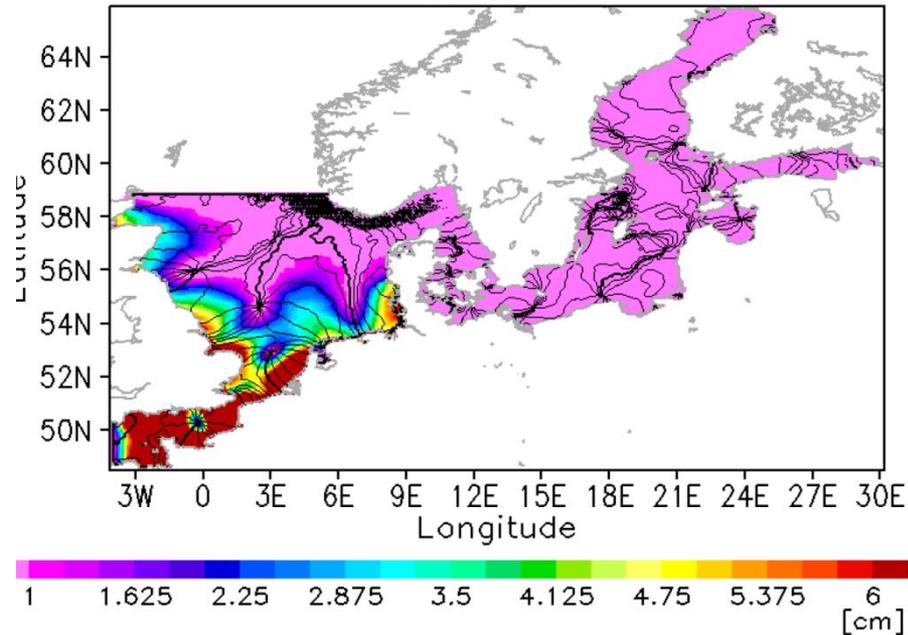
AMM7 M4



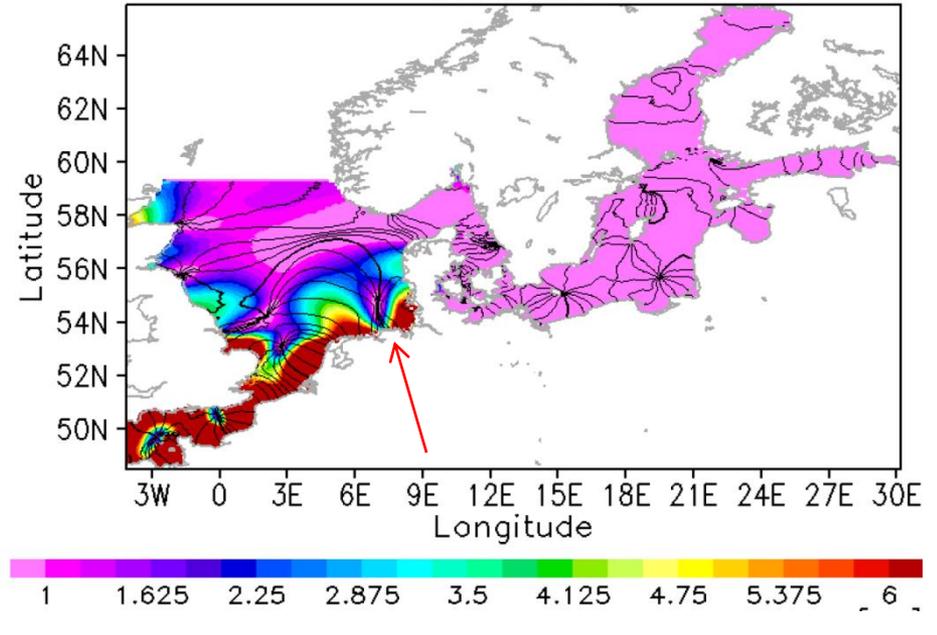
CMOD4 M4



GETM M4



NEMO M4

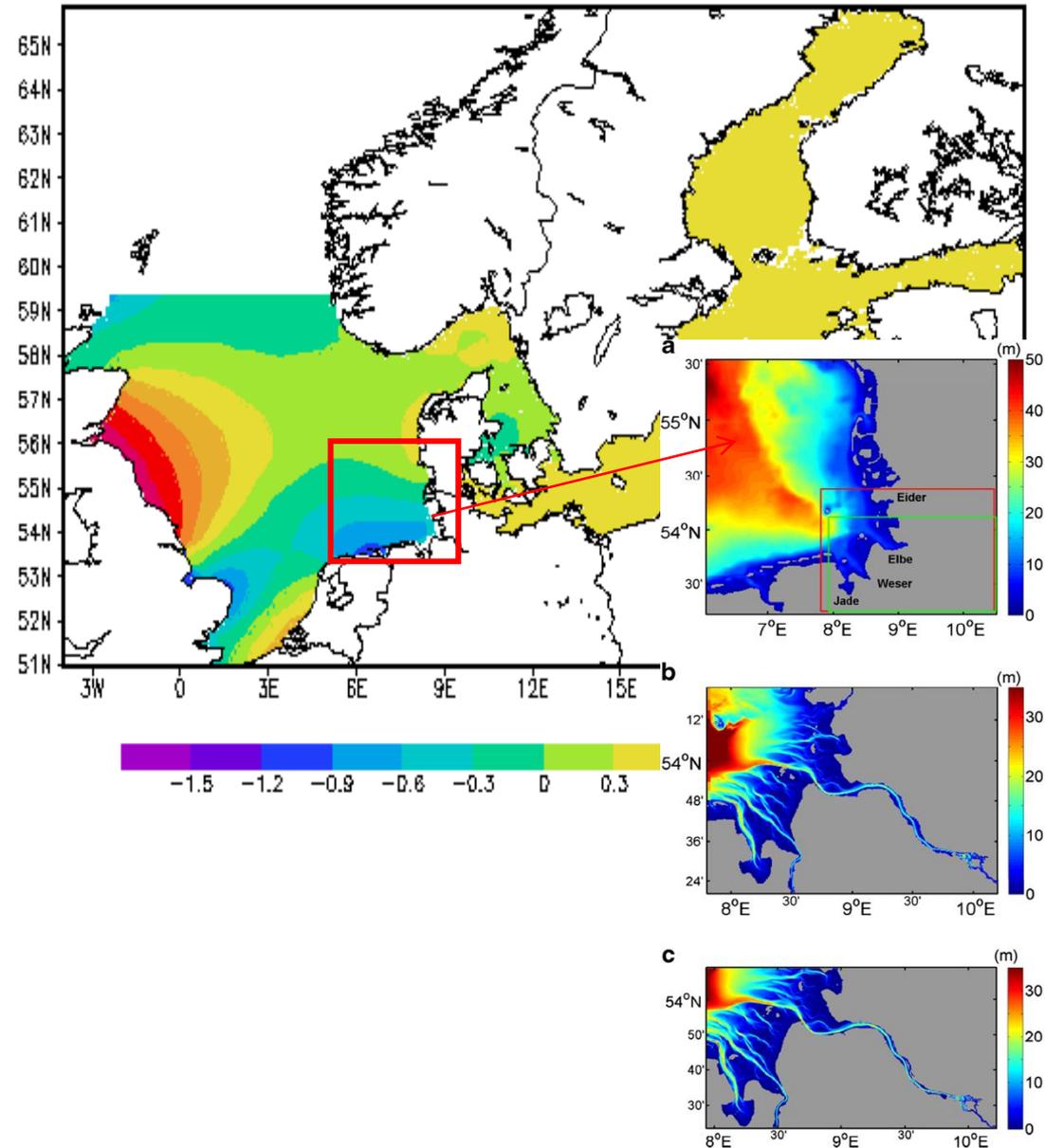


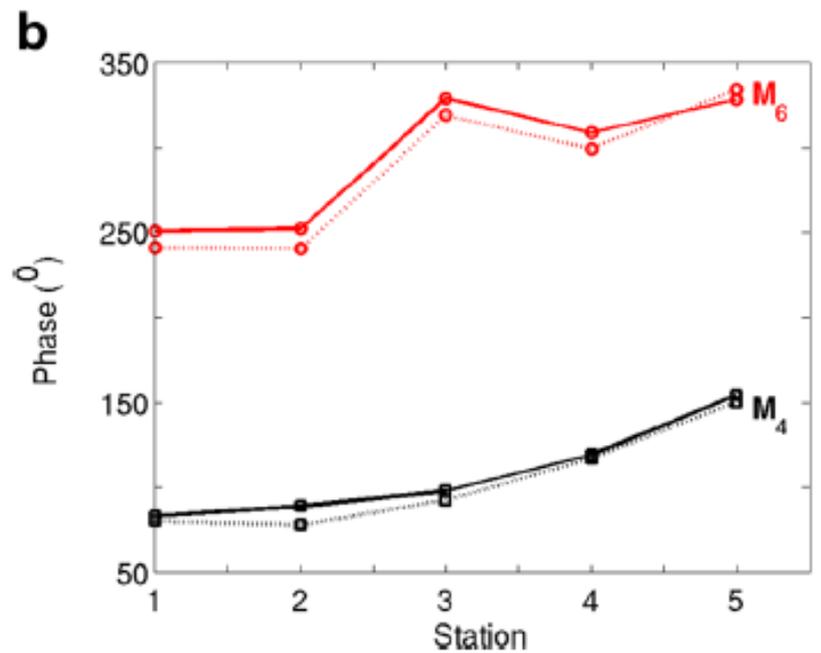
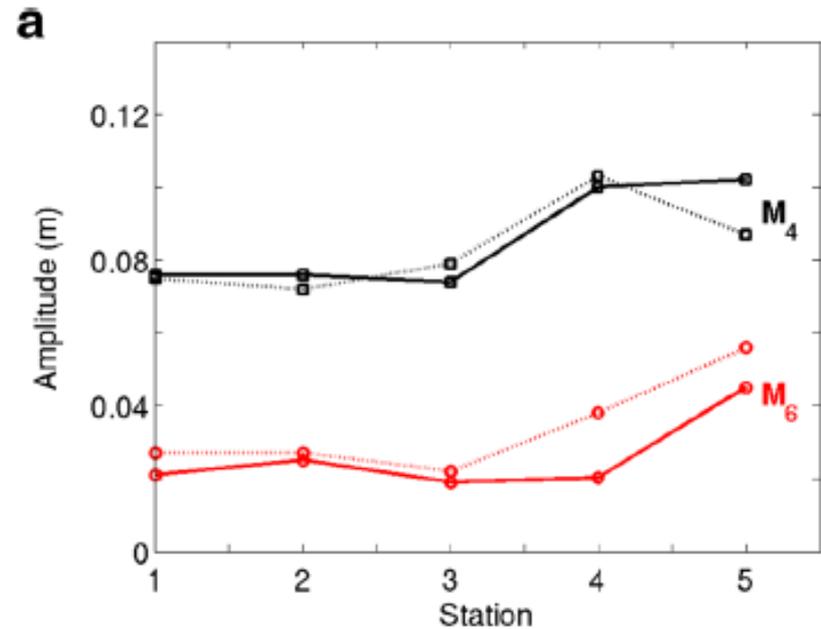
Nested Modeling System: GETM

The modeling system consists of three nested models:

- ***North Sea-Baltic Sea model (5 km)***
- ***German Bight model (0.8 km)***
- ***Wadden Sea model (200 m)***
- ***Forcing***
 - ***Atmospheric forcing (DWD 1-hr atmospheric forecasts),***
 - ***River run-off – hourly data,***
 - ***Open BC – tides, T and S***

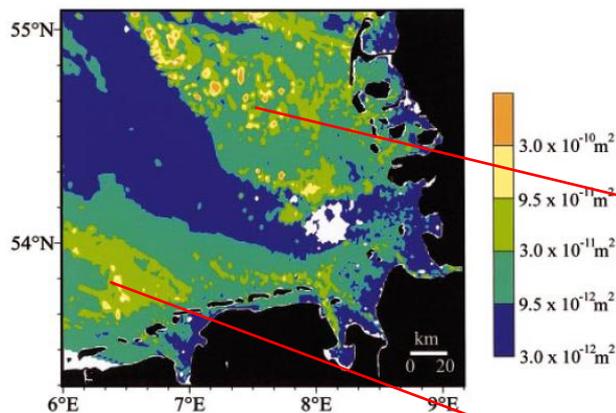
Surface Elevation (meters)



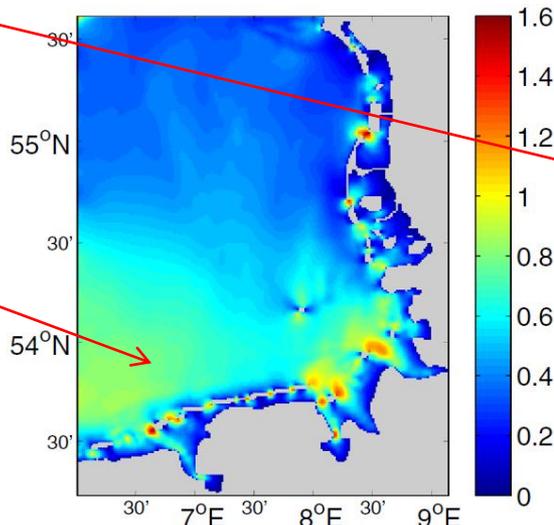
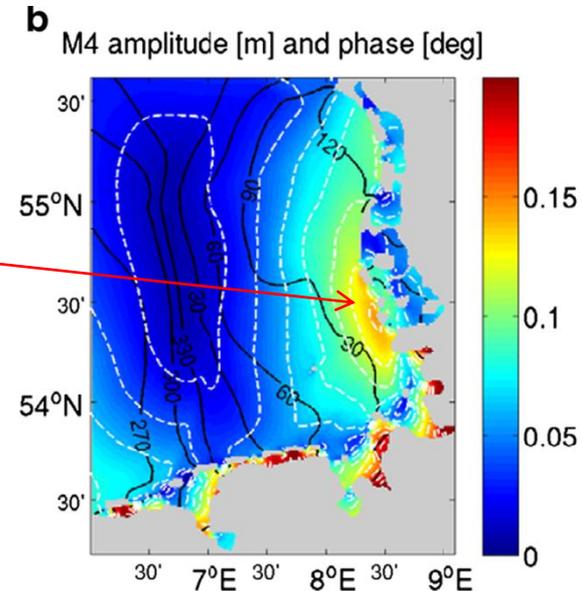
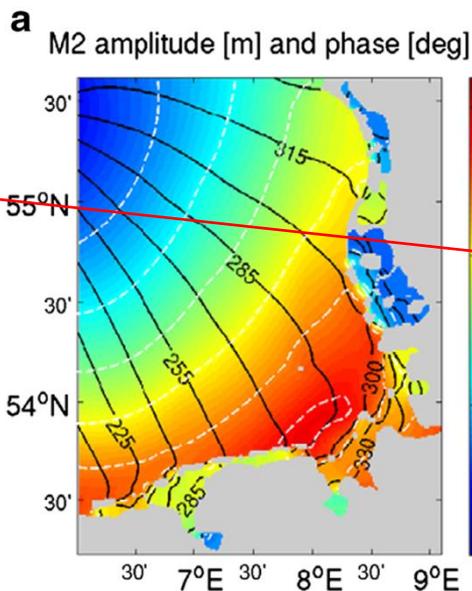


Validation: Stanev et al. (2014), ODYN, Observations-full line

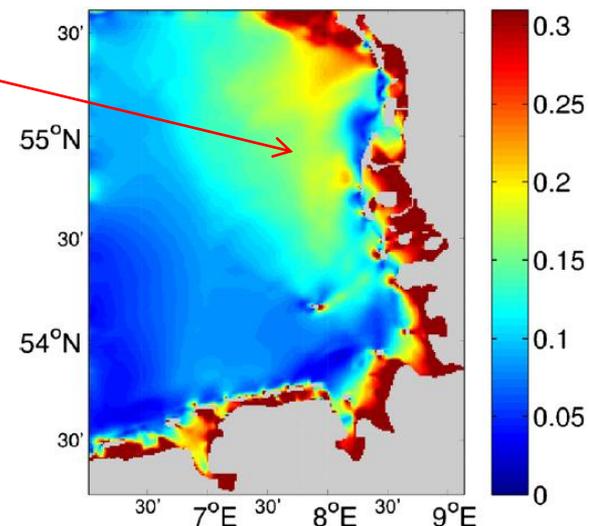
The region of strongest generation of overtones (reflection and refraction of a Kelvin wave)



Janssen et al., 2005



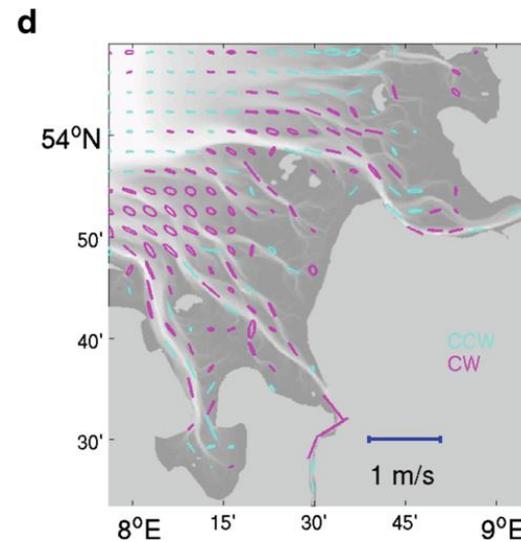
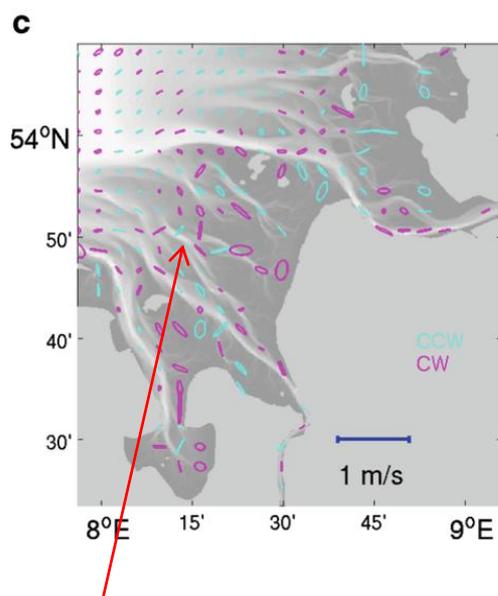
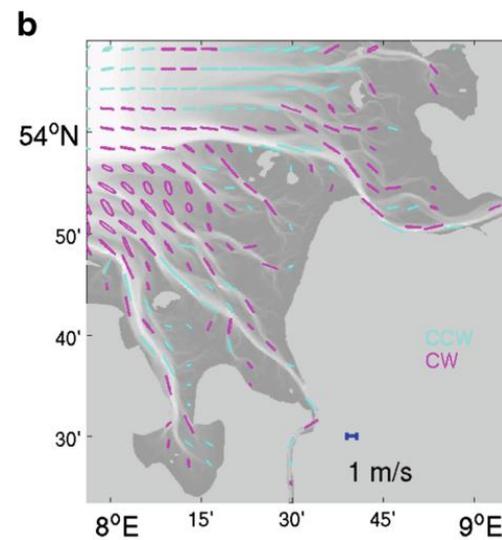
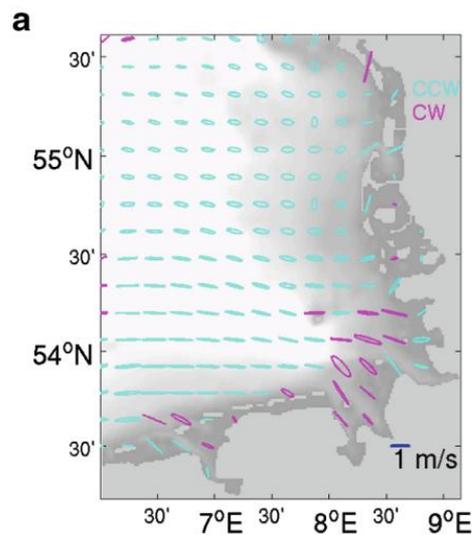
Amplitudes of the major axes currents (M2)



M4/M2 (tidal distortion) roughly represents the effect of energy transfer from M2 to M4

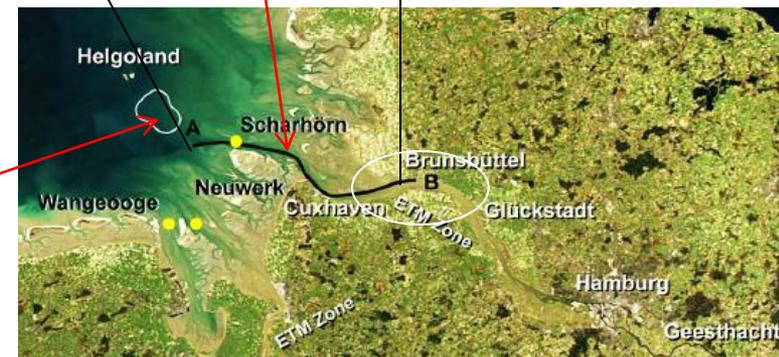
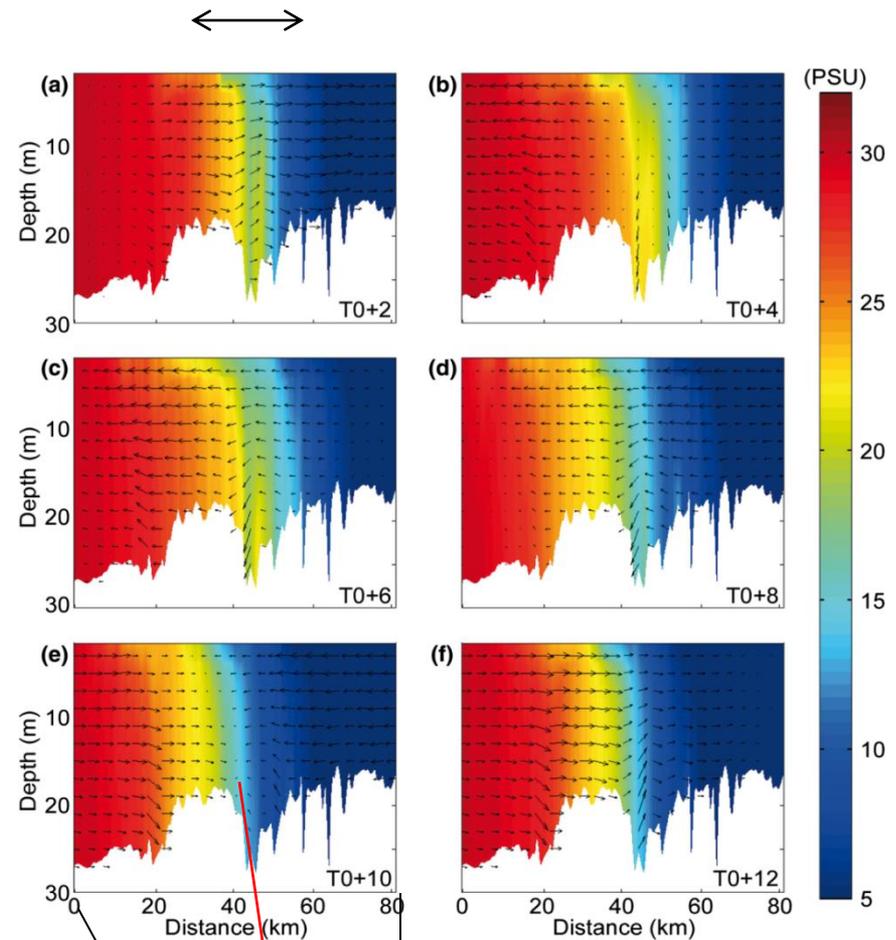
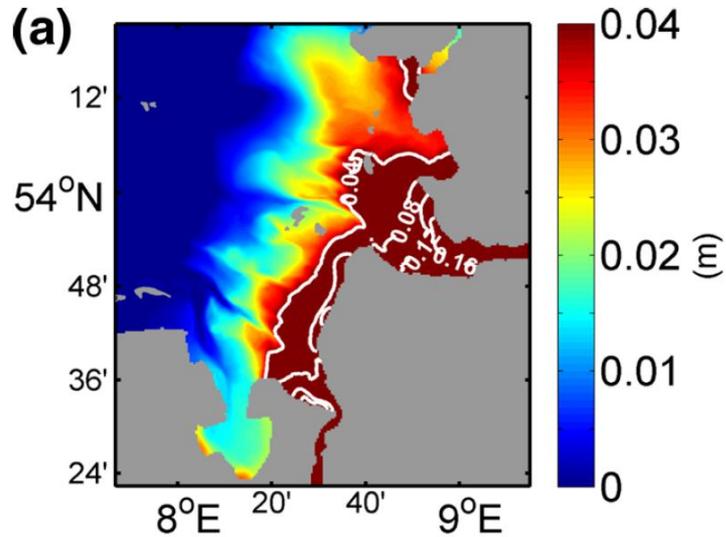
Conclusion:
It is not only the energy, but also the tidal distortion which shapes the governing processes (tidal asymmetry)

The control of bathymetry



higher harmonics reveal “orthogonal” patterns, which could drive secondary circulation

The role of baroclinicity (salinity front/



Roughly the net sediment transport is ebb- or flood dominated in cases of short ebb or flood, respectively

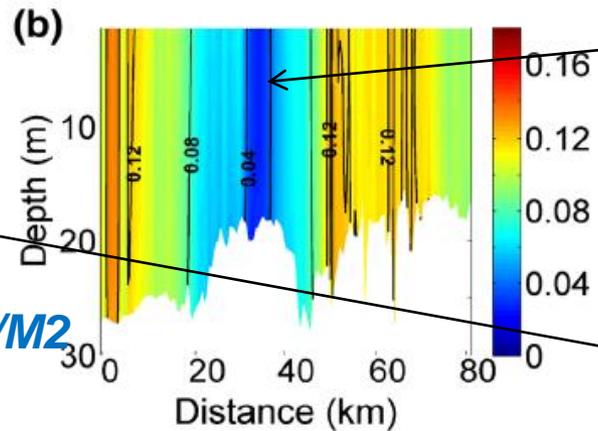
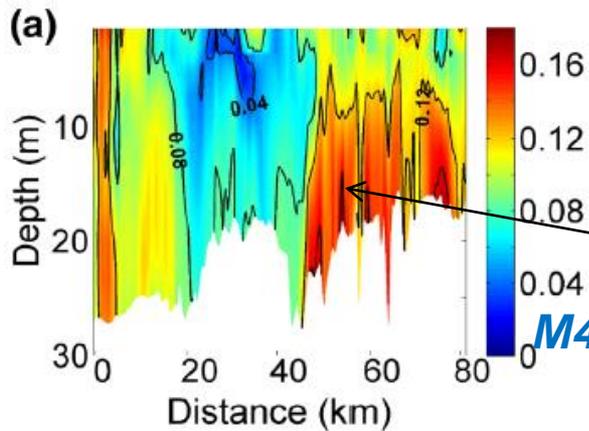
Zone of accumulation of sediment

baroclinic

barotropic

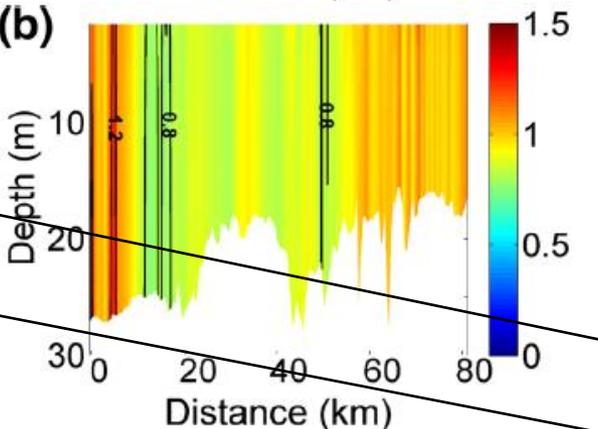
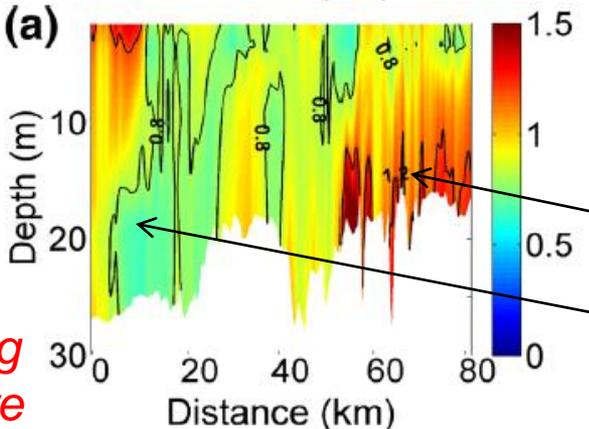
*M4/M2,
energy transfer*

spring



*bottom intensified
ETM
(flood dominated)*

spring



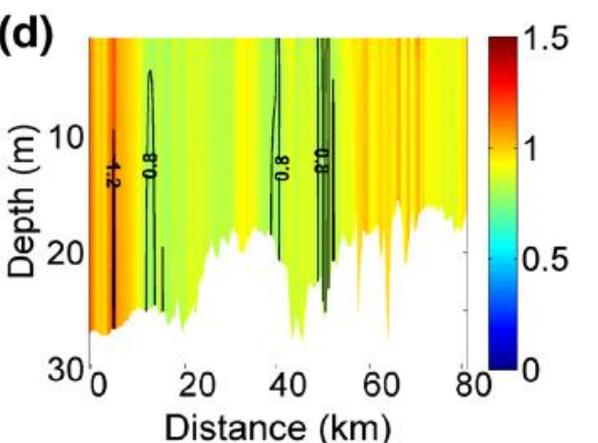
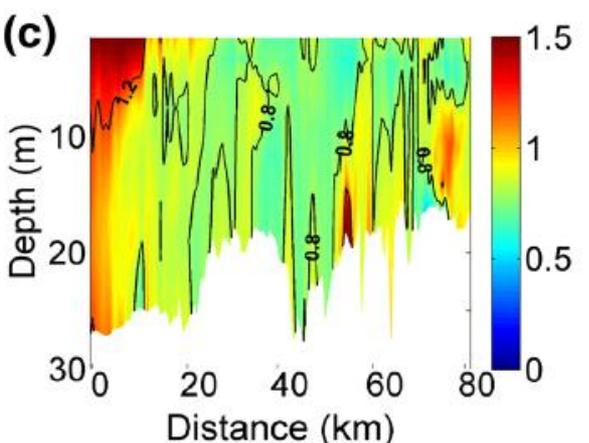
*Maximum flood

maximum ebb

flood versus
ebb dominance*

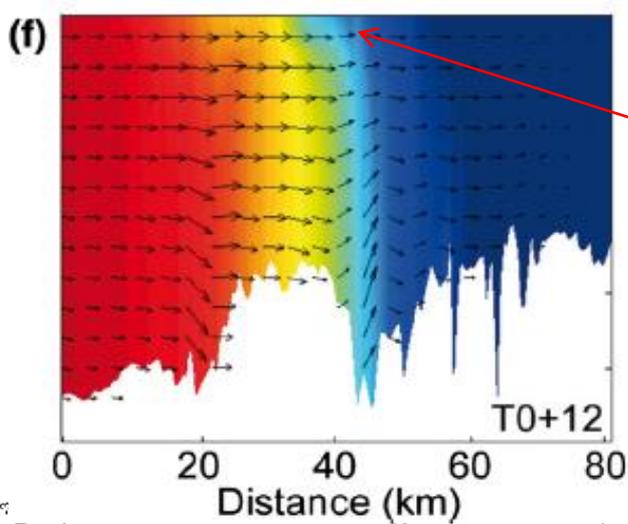
*Neap-to-spring
differences are
enhanced by
density.*

neap

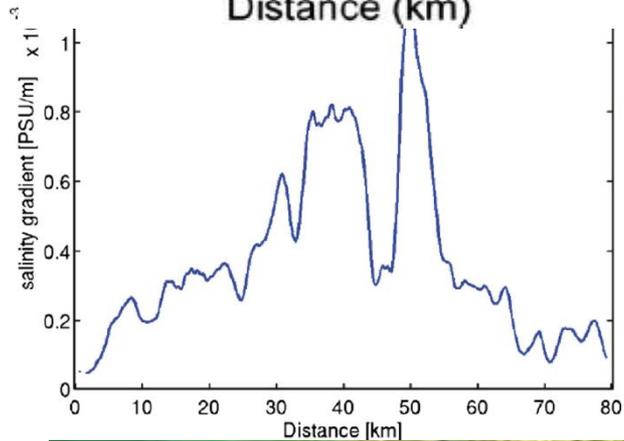


*Maximum flood

maximum ebb*



gravitational
circulation



The *non-linearities* and the resulting distortion in the tidal signal occur in the area of the *salinity front*.

The largest tidal *excursions* are associated with the maximum *divergence/convergence* of currents in the mouth.

$$Si = \frac{b_x H^2}{C_D U_T^2}$$

competition
between
baroclinic and
barotropic
forcing

$$L_T = U_T T_T / \pi$$

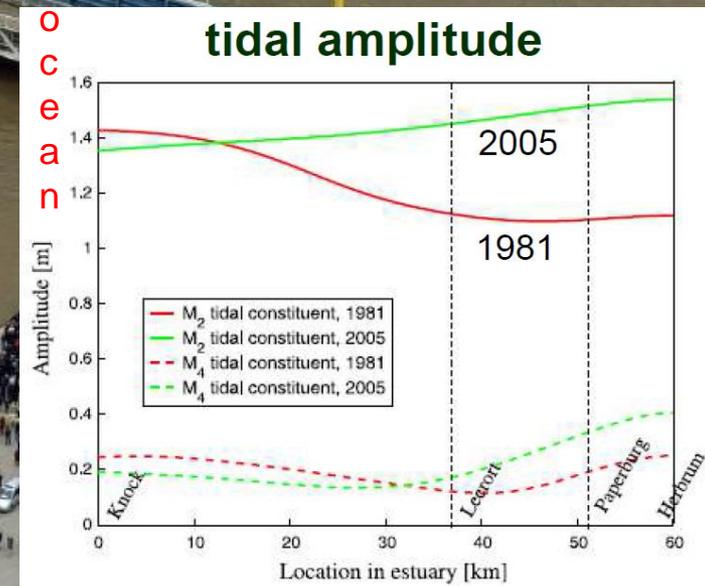
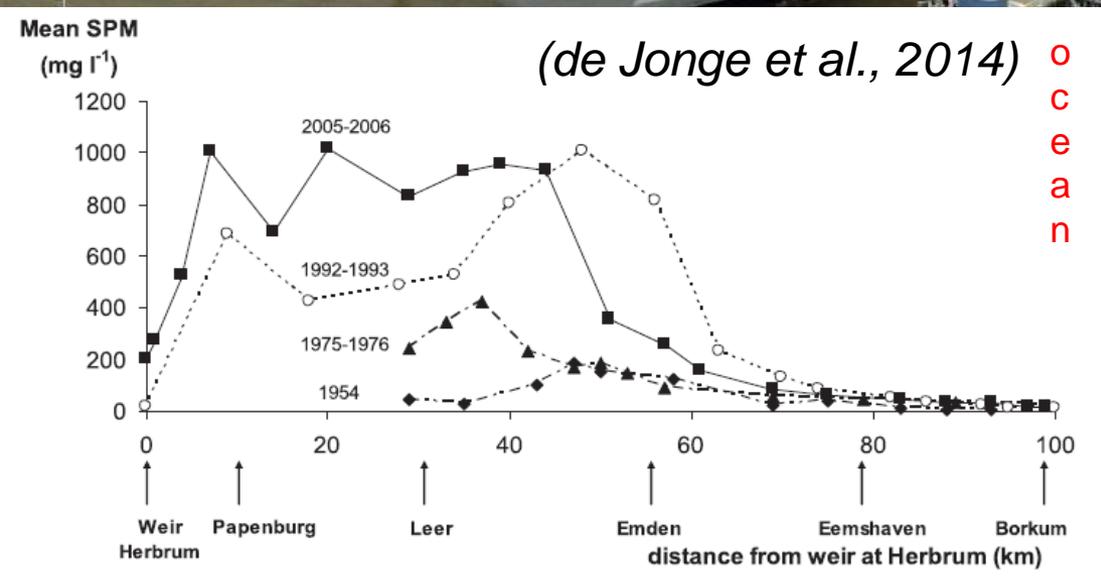
The correlation between salinity and velocity tended to enhance the *vertical overturning*.

Further estuarine applications

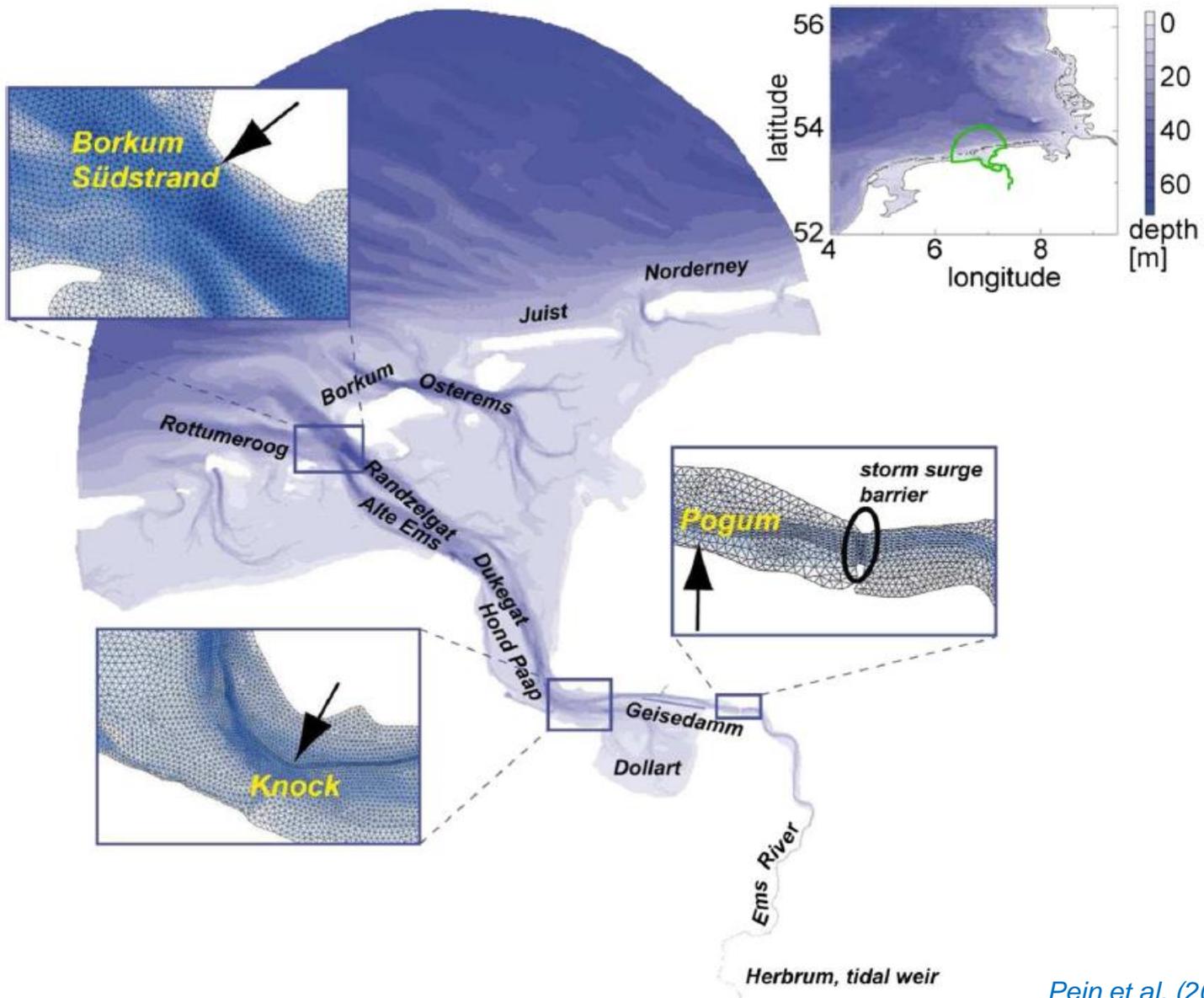
Welcome to Ems



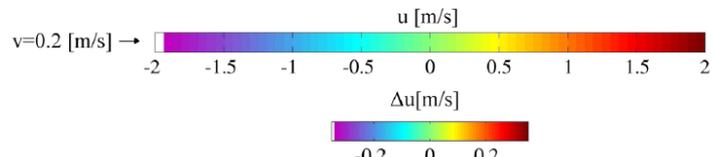
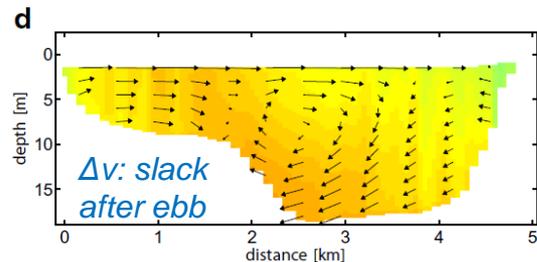
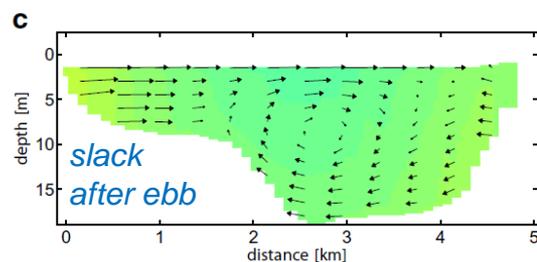
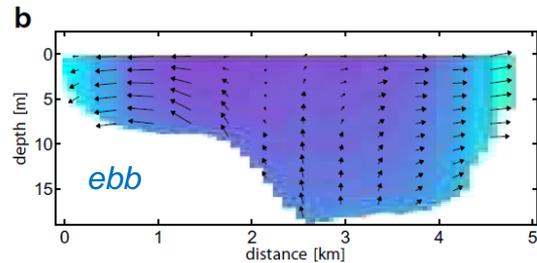
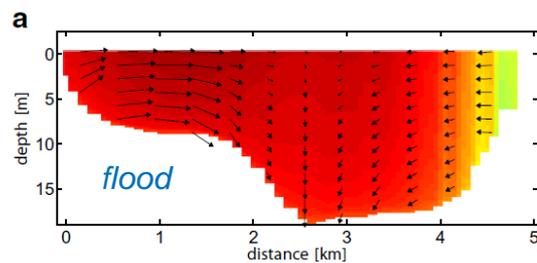
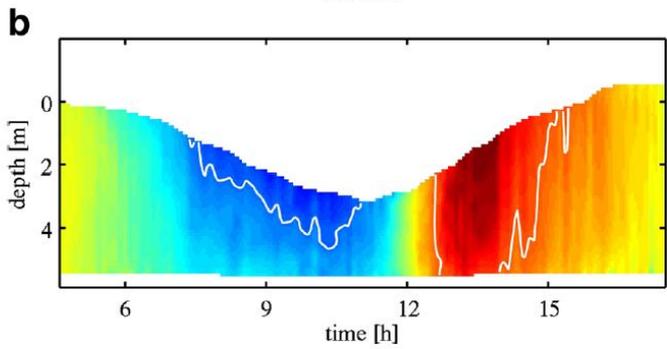
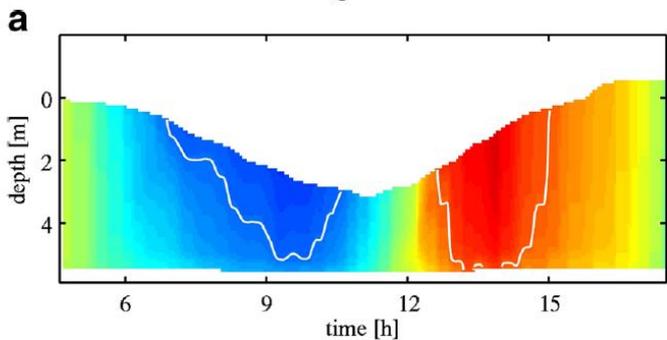
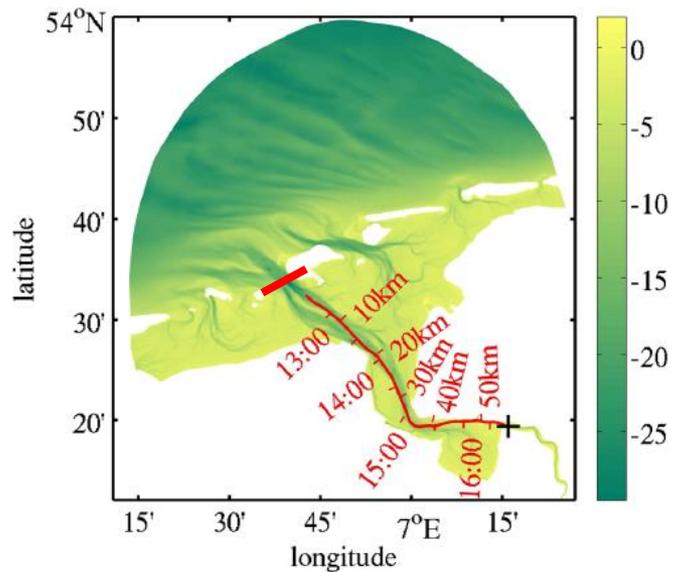
Schuttelaars et al., 2012



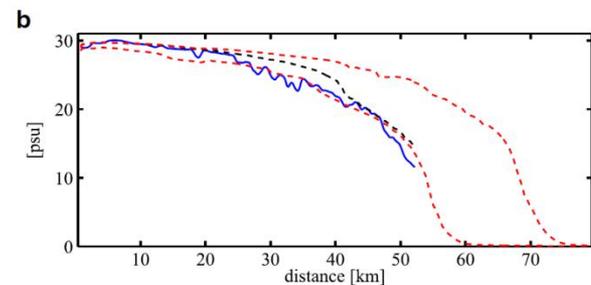
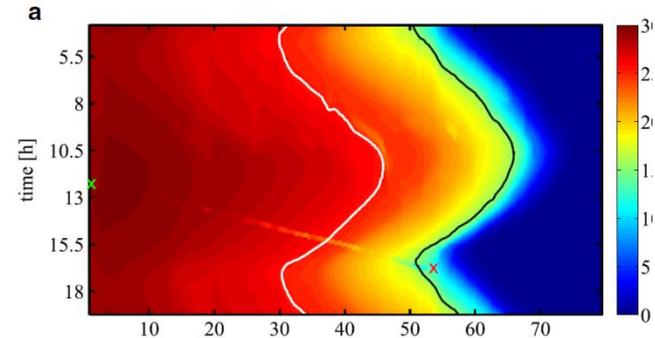
Semi-implicit
Cross-scale
Hydroscience
Integrated
System Model
(SHISM/SELFE)



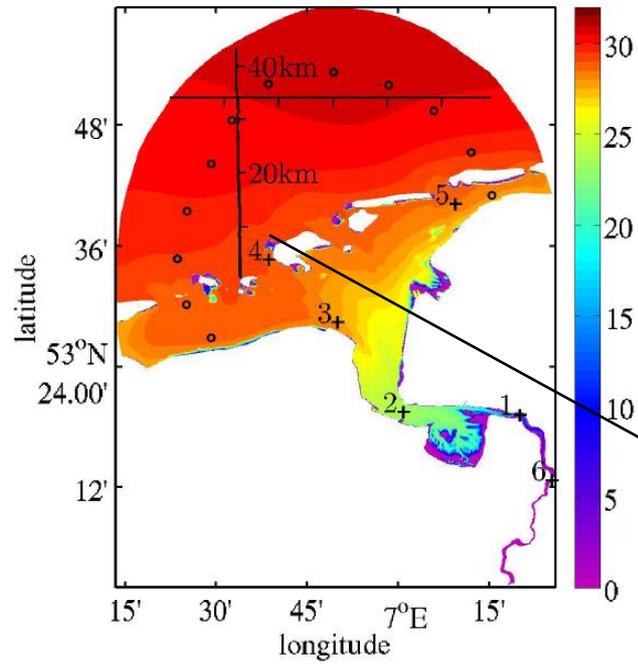
Validations



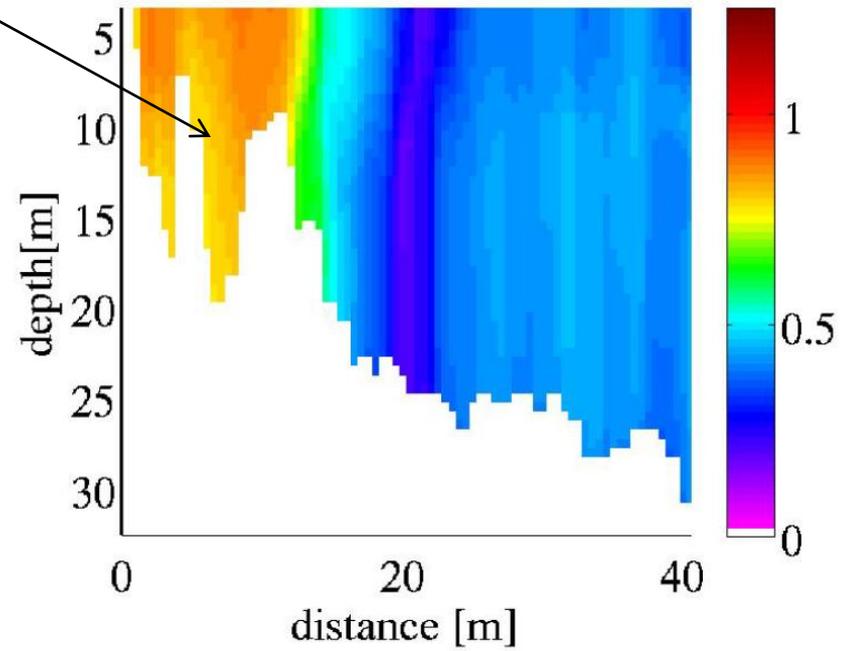
Non-canonical secondary circulation



Salinity



M4 Salinity



Conclusions:

- *The present research: an illustration of the **interaction between the tidal flow and the density gradient** in a region influenced by fresh water.*
- ***Estuarine theories** should probably account for the joint contributions of nonlinearities and baroclinicity.*
- *The vertical patterns of tidal asymmetries and their possible role for **material transport** and deposition deserve deeper consideration*